Uptake and Storage of CO₂ in Subtropical Mode Water (STMW) of the North Atlantic Ocean

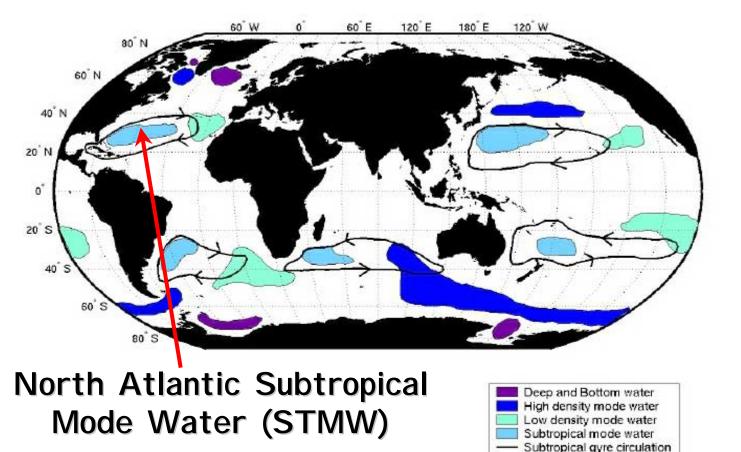
Nicholas R. Bates Bermuda Biological Station For Research (BBSR)





STMW Carbon Uptake and Storage

• The magnitude and interannual variability of uptake and storage of carbon dioxide (CO_2) and storage into mode waters are poorly quantified.



Source: Talley, 2000

STMW in the North Atlantic Ocean

4°C

18°C

Gulf Stream

Generic winter location of STMW formation

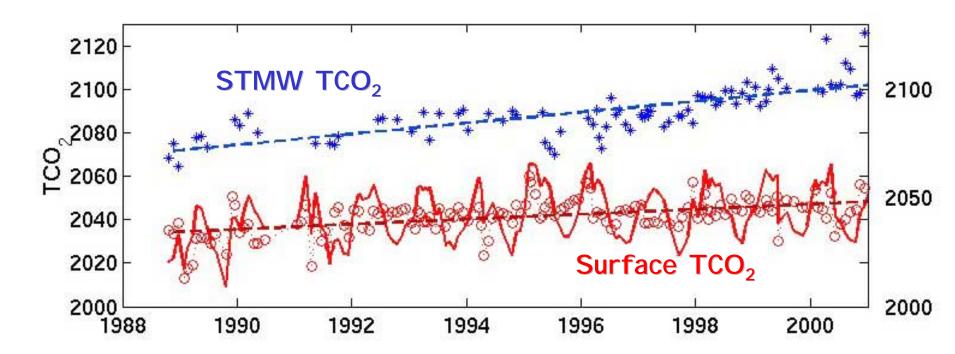
18°C 25°C

Geostrophic recirculation pathways of STMW

Interannual variability of STMW formation is primarily associated with climate variability (i.e. North Atlantic Oscillation, NAO)

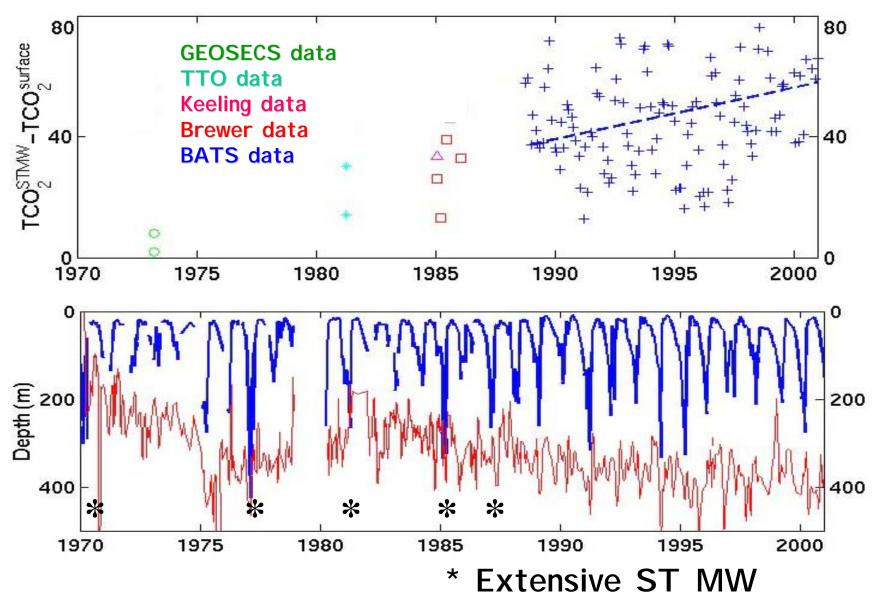
February SST composite image source: Norm Nelson, UCSB

Interannual Variability of CO₂



• Mixed layer TCO_2 (µmoles kg⁻¹ yr⁻¹) TCO_2 +1.20 + 0.35 (r²= 0.36*) $nTCO_2$ +1.19 + 0.25 (r²= 0.37*) DO -0.10 + 0.24 (r²= 0.00) • STMW TCO_2 (µmoles kg⁻¹ yr⁻¹) $TCO_2/nTCO_2$ +2.22 + 0.27 (r²= 0.65) DO -0.58 + 0.22 (r²= 0.27)

Non-steady state changes in CO₂

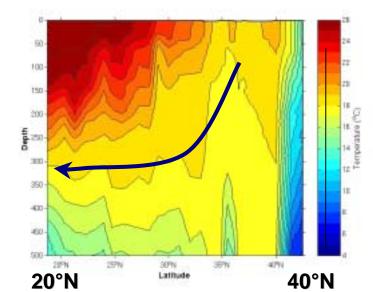


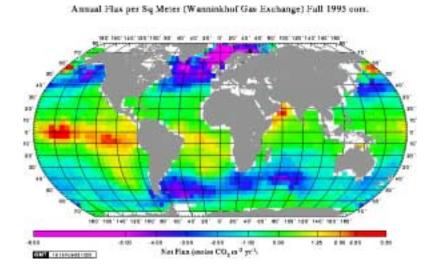
Increased Oceanic CO₂ Sink in 1990's

- Annual global ocean uptake of CO_2 is about 2 Pg C yr⁻¹.
- Over the last 12 years, the extra uptake of CO_2 into STMW (~ 0.6 to 2.8 Pg C) has the same range as the global annual uptake of CO_2 .

• Since 1988, STMW has taken up (~ 0.05 to 0.23 Pg C yr⁻¹). This is ~3 to 11% of the best estimate of annual uptake of CO_2 into the global ocean.

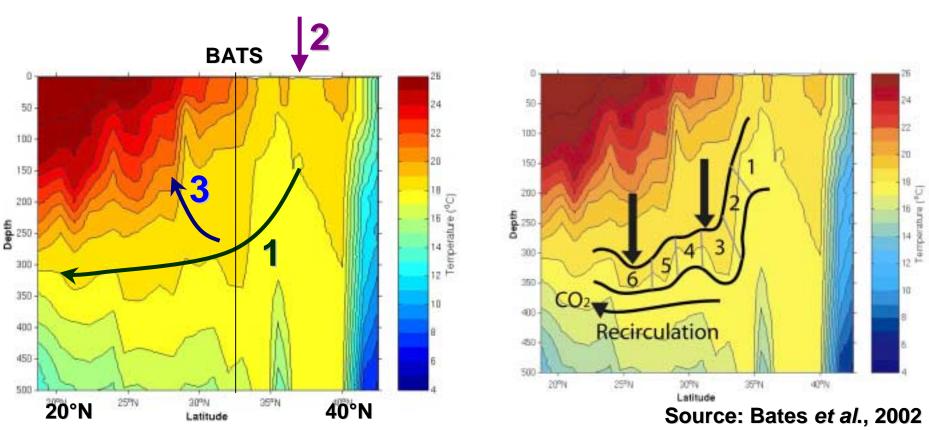
3-11% of global CO_2 uptake





Causes for changes in CO₂

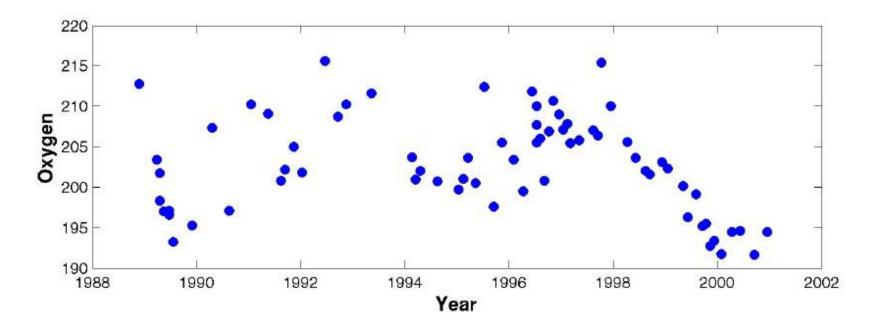
- 1. Remineralization of OM (sampling older water over time)
- 2. Changing flux of CO₂ through gas exchange
- 3. Retention of CO_2 or loss from STMW (by mixing)



Increase in Remineralization Rate?

DO $-0.58 \pm 0.22 \ \mu \text{moles kg}^{-1} \ \text{yr}^{-1} (r^2 = 0.27)$ Nitrate $-0.02 \pm 0.02 \ \mu \text{moles kg}^{-1} \ \text{yr}^{-1} (r^2 = 0.15)$ Phosphate $-0.00 \pm 0.00 \ \mu \text{moles kg}^{-1} \ \text{yr}^{-1} (r^2 = 0.13)$ [Temperature $+0.003 \pm 0.004 \ ^\circ\text{C} \ \text{yr}^{-1}$ Salinity $+0.002 \pm 0.000$ yr⁻¹]

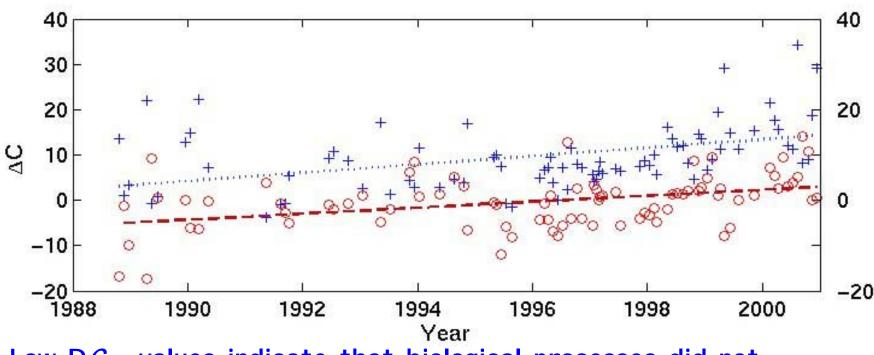
• STMW TCO₂ changes <u>not</u> due to remineralization (i.e., decrease in DO) or sampling of older water.



Increase in Remineralization Rate?

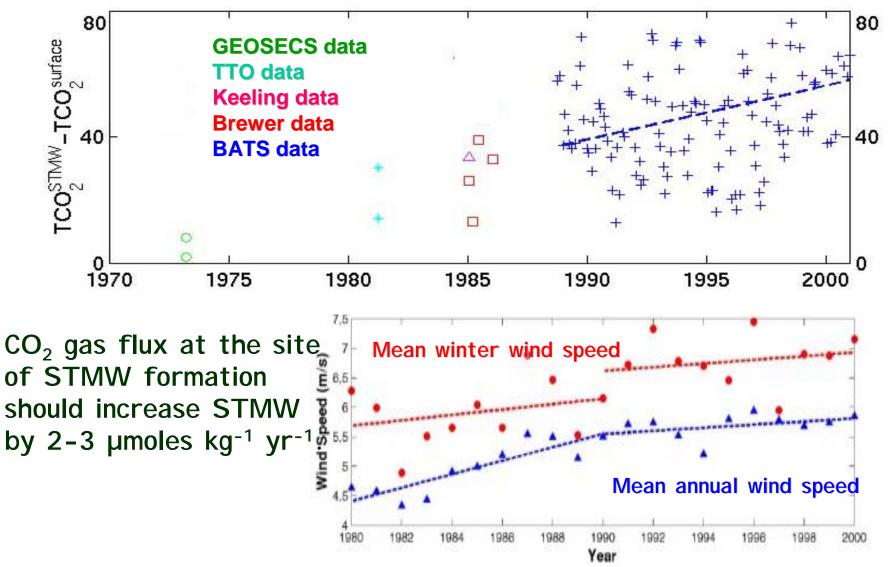


+2.22 (μmoles kg⁻¹ yr ⁻¹) +0.90 (μmoles kg⁻¹ yr ⁻¹) +1.19 <u>+</u> 0.26 (μmoles kg⁻¹ yr ⁻¹) (r²= 0.47) +0.28 + 0.12 (μmoles kg⁻¹ yr ⁻¹) (r²= 0.25)



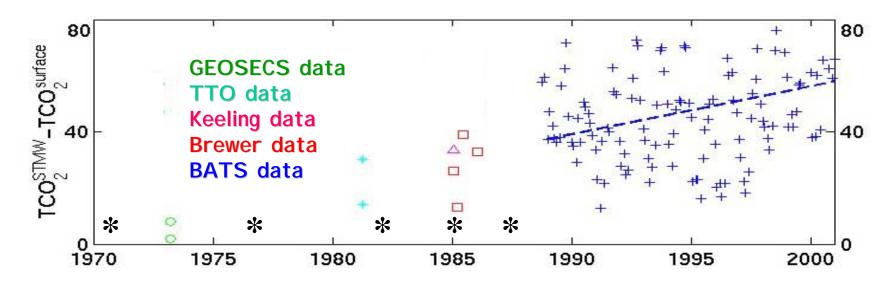
• Low DC_{bio} values indicate that biological processes did not contribute much to the +2.2 µmoles kg⁻¹ yr ⁻¹ change in STMW TCO₂ Source: Bates *et al.*, 2002

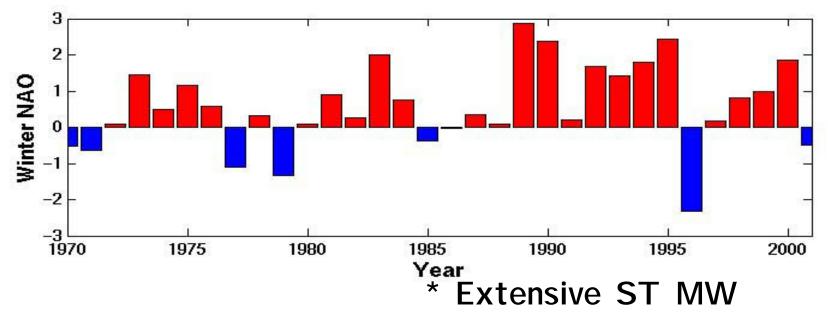
Increased Gas Exchange?



Higher winter wind speeds in 1990's compared to 1980's

Variability of Carbon Storage?

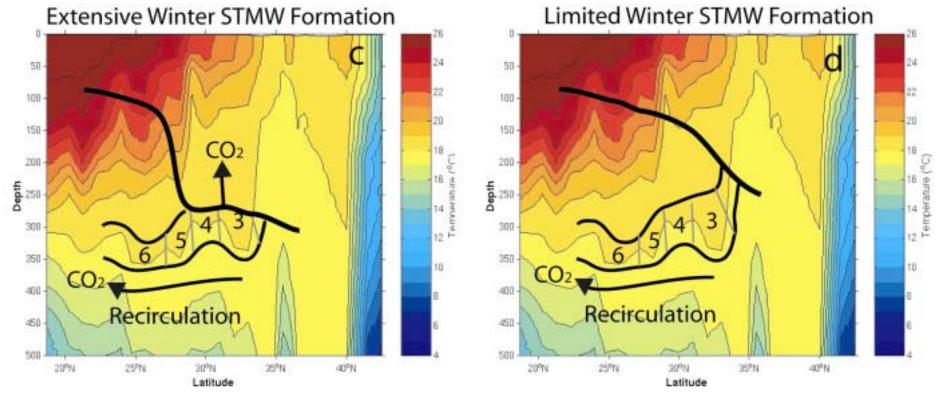




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NAO-ve State NAO+ve State

1960's, 1970's, early 1980's: Post 1987: CO_2 transferred CO_2 in STMW redistributedto ocean interior



Short-term CO₂ sink ~1-4 years

Long-term CO₂ sink >10 years

Conclusions:

•Apparent coupling between modes of climate variability such as NAO and CO₂ uptake and storage in the subtropical gyre.

• Since 1988, $\sim 0.6-2.8$ Pg (10¹⁵ g) of extra CO₂ has accumulated within the gyre STMW layer. This represents a longer term oceanic sink of CO₂ (>10 years) in the 1990's compared to earlier decades, and <u>3-11%</u> of global ocean flux?



I mage Source: BBSR

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ERE DISCOVERIES BEGIN